

1 **Aquifer contamination by coastal floods in the plain of**
 2 **Costa da Caparica, Almada, Portugal**

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8 **Abstract.**

9 Climate change might lead to sea level rise, changes in the frequency, intensity and
 10 duration of wave storms, which will lead to an increase of coastal flooding events.
 11 One of the most important long-term effect of coastal floods is saltwater intrusion
 12 induced by the vertical infiltration of the salt water behind the overtopped and/or
 13 breached coastal barriers. A single overflow event may contaminate the freshwater
 14 aquifer for several years until it is remediated naturally under the effect of precipita-
 15 tion and subsequent seaward directed flow.

16 The main objective of this study is to understand the effects of a maritime storm
 17 and the induced hinterland inundation on the water quality of a coastal aquifer. The
 18 study targets the municipality of Almada (Portugal) which has an extensive coastline,
 19 with coastal aquifers with high susceptibility to contamination, namely the unconfined
 20 aquifer of the coastal plain of Costa da Caparica.

21 Groundwater flow modelling and mass transport in the aquifer were modelled us-
 22 ing the MODFLOW and MT3DMS software. The aquifer contamination by an over-
 23 flow was modelled considering wave overtopping and flooded area on the coast was
 24 estimated from previous works that coupled a wave transformation model (SWAN)
 25 with a high-resolution swash model (XBEACH). The extent of the subsurface con-
 26 tamination is a function of the flood extent.

27 Results showed that at an extreme storm event, waves can overtop the coastal dune
 28 causing a coastal flood that extends approximately 160 meters inland with a signifi-
 29 cant increase in chlorine concentration in the aquifer. Recover of the aquifer to previ-
 30 ous concentrations was found to take several years.

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33 **Keywords:** Coastal Flood, Coastal Aquifer, Salinization, Numerical Modeling,
 34 Mass Transport.

35 **1 Introduction**

36 Coastal aquifers are highly susceptible to several natural and anthropic related en-
 37 vironmental changes such as sea level rise, reduction of recharge, pollution or coastal
 38 floods by wave overwash. These latter processes can induce significant saltwater

39 intrusion in freshwater aquifers (Elsayed & Oumeraci, 2017). Contamination of
40 groundwater by infiltration of seawater is a consequence that deserves attention.

41 In the coastal region of the municipality of Almada, there is an unconfined aquifer
42 developed in sandy sediments that is exploited for irrigation purposes. This aquifer
43 presents high susceptibility to contamination, especially in relation to the advance of
44 the salt wedge, caused by sea level rise and overexploitation. Coastal flooding events
45 can also have a significative impact, because they are fast and cover vast areas of
46 infiltration. The salinization of the aquifer can undermine local economic activity
47 linked to agriculture and the exploitation of groundwater in shallow wells.

48 The main objective of this study is to understand the effects of a extreme maritime
49 storm and the associated hinterland inundation on the water quality of a coastal aquifer.
50 The study targets the municipality of Almada (Portugal).

51 The municipality of Almada is located on the central coast of mainland Portugal,
52 on the left margin of the Tagus River (NW region of the Setúbal Peninsula), Setúbal
53 district. The plain of Costa da Caparica is located on the NW extreme of the municipality.
54 It is limited by the Tagus River (and its estuary) at the north and at west by the
55 Atlantic Ocean, and by a cliff at east (Figure 1).

56 The regional climate is marked by an oceanic regime with warm winters and hot
57 summers, by proximity to the Atlantic Ocean (and whose weak annual thermal amplitude
58 allows a regularization of temperature throughout the year). The annual average
59 temperature is between 16 and 17.5 °C, and the average annual precipitation between
60 500 and 700 mm (in Caria *et al.*, 2013). During the winter period the plain of Costa da
61 Caparica is recurrently affected by wave overtopping, with consequent flooding of
62 inland areas.

63 The plain of Costa da Caparica is formed by deposits composed of dune sands,
64 slope deposits and alluviums and / or landfills of Pliocene - Holocene (Pais *et al.*,
65 2006). Those deposits cover sedimentary formations of the Miocene also represented
66 on the eastern cliff. The Miocene formations correspond to continental deposits alternated
67 by other marine ones, characteristic of a broad alluvial plain in the form of an
68 estuary, subject to transgressions.

69 The aquifer developed in Pliocene - Holocene deposits is unconfined, with thickness
70 between 20 m – 25 m. The hydraulic conductivity of the formations varies from
71 3.4 m/day to 4.5 m/day, and Ferreira (2012) estimated an average value of 704 m²/day
72 for the transmissivity. The aquifer receives direct recharge from precipitation, around
73 250 mm per year, and lateral recharge from the Miocene formations at east (Ferreira,
74 2012) but the amount of the water transfer is not well known.

75 **2 Materials and Methods**

76 To achieve the objectives, the groundwater level and flow in the coastal aquifer
77 were modelled using the MODFLOW software. The mass (chloride concentration)
78 transport using the MT3D allowed to understand the aquifer contamination associated
79 with floods during storm events. Both softwares are compiled by the Processing

80 Modflow (PMWIN) software (Version 5.3.1, by W. H. Chiang and W. Kinzelbach,
81 2001).

82 The aquifer was simulated in 3D and at steady-state regime, using a difference el-
83 element mesh and one unconfined layer with 20 m of thickness. The mesh is composed
84 by 440 columns and 430 lines with variable cells width (10, 5 and 2.5 meter) with
85 higher resolution in the area affected by floods. The boundary conditions define the
86 groundwater influx from the east, from the Miocene aquifers, and the constant head at
87 sea side. The hydraulic conductivity of the layer was obtained by interpolation (in-
88 verse distance method) of values obtained through granulometric analysis of selected
89 samples (equivalent diameter) and varies between 0.035 and 0.028 m/s (Figure 1).
90 Considering the sandy composition of the plain aquifer a porosity value of 0.25 was
91 assumed. The recharge by precipitation was obtained from Caria (2012, 2013) and
92 ranges from 0 to 254 mm per year (Figure 1). The zero represents the impermeable
93 zones (anthropic uses and actions), while the higher permeable zones are associated to
94 agricultural land. The initial hydraulic heads, interpolated by kriging, are represented
95 in Figure 1; the water level varies from 4.43 m at east to close to 0 m (mean sea lev-
96 el).

97 The mass transport model was simulated in transient regime for a particular section
98 of the model (location in Figure 1), representing the area flooded during the storm on
99 8 - 10 of February of 2014. The highest volume of flood was estimated in 0.2 L/m/s,
100 coupling a wave transformation model (SWAN) with a high-resolution swash model
101 (XBEACH), (Pires, 2017). The extension of the flood was estimated in approximately
102 162 m inland, considering a wave overtopping of 600 seconds. The initial conditions
103 for the chloride concentration were assumed as 60 mg/L for the regional groundwater,
104 5 mg/L for rain water and 20,000 mg/L for sea water. The simulation was carried out
105 in two steps: the first consider the infiltration of the seawater on the top of the layer,
106 representing a flood of 1 day; the second is equivalent to the period of natural remedi-
107 ation, and the chloride concentration was simulated for different periods (stress peri-
108 ods), 1, 2 and 5 years), to understand how long the aquifer will takes to recover.
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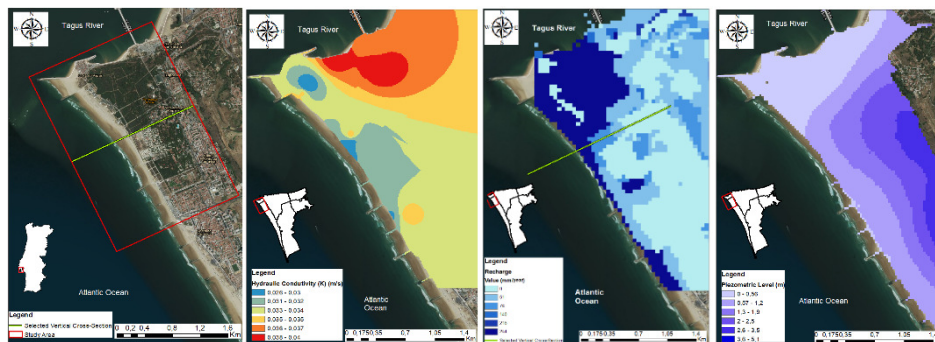
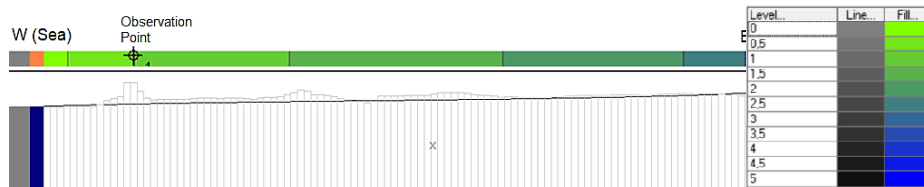


Figure 1 – Location and limits of the study area; spatial distribution of the hydraulic conductivity; variation of the direct recharge of the aquifer; piezometric map of the study area; (from left to right).

111 3 Results and Discussion

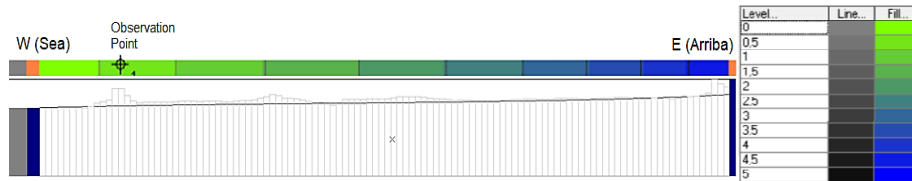
112 3.1 Groundwater flow modelling

113 The aquifer numerical modeling show that the groundwater flows from east to west
 114 into the sea, corroborating the results of Ferreira (2012) and Storm (2014). Figure 2
 115 show a cross section of the simulated groundwater level in the aquifer, at steady-state
 116 conditions. It is possible to see that the infiltration of the seawater in the inundated
 117 area causes a rise in the groundwater level, that can be higher than the beach profile.
 118 In this simulation the piezometric level varies between 5 m at east and 0.16 m at the
 119 beach. The second simulation shows (Figure 3) that the heads recover when the storm
 120 ends and return to the normal values, i.e., 0 m close to the sea.



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122 *Figure 2 - Vertical section of the aquifer with the simulated piezometric level (dark line), when wave*
 123 *overtopping occurs on the study area (stress period 1).*



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125 *Figure 3 - Vertical section of the aquifer with the simulated piezometric level (dark line), after the storm*
 126 *(stress period 2).*

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128 3.2 Mass Transport Modelling

129 The mass transport modelling results, in transient regime and two steps, are repre-
 130 sented in Figures 4 and 5. The first simulation corresponds to the flood period, with
 131 an inundation period of 1 day; the second concerns to the evaluation of the time the
 132 aquifer needs to recover from seawater contamination during the flood.

133 At an observation point in the aquifer (Φ), located at the limit of the flooded area,
 134 it is possible to see that the concentration simulated in the aquifer after 1 day of flood
 135 is 19,177 mg/L (stress period 1; Figure 4).

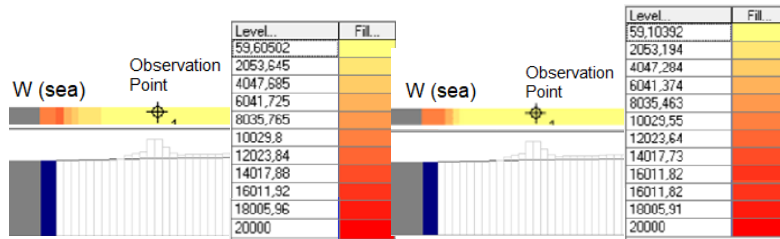
136 The evolution of the chloride concentration in the aquifer over time, at the observa-
 137 tion point, can be seen in Figures 5 and 6 where is clear that the aquifer needs several
 138 years to recover the groundwater quality. At the time steps length of 1 to 5 years, the
 139 simulated chloride concentration at the observation point will be 318, 84 and 67
 140 mg/L, respectively.

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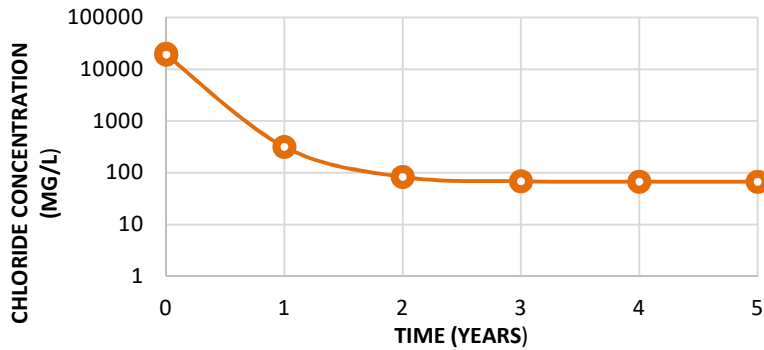
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144 *Figure 4 – Vertical section of the aquifer with the simulated chloride concentration in observation point after 1 day of flood; square represents the aquifer zone in Figure 5; dark line is the groundwater level.*

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149 *Figure 5 - Vertical section of the aquifer zone, delimited in Figure 4, with the simulated chloride concentration after 1 year, 2 years and 5 years of the storm; dark line is the groundwater level.*

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153 *Figure 6 - Chloride concentration calculated on the vertical cross- section, after the wave overtopping occurs on the study area, versus time (years).*

154 **4 Concluding Remarks**

155 The numerical simulation of the flow model and mass transport of the unconfined
156 aquifer of the Costa da Caparica Plain shows that the coastal flood associated with
157 extreme events can lead to the contamination of the aquifer over a long period of
158 time.

159 During a storm, a wave with a flow rate of 0.2 L/m/s can cause a wave overtopping
160 that reaches approximately 160 meters in land. The infiltrated seawater can contami-
161 nate the aquifer for a period that can reach several years.

162 **Acknowledgements**

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